



Discovery and significance of Cretaceous fossils from the Xigaze Forearc Basin, Tibet

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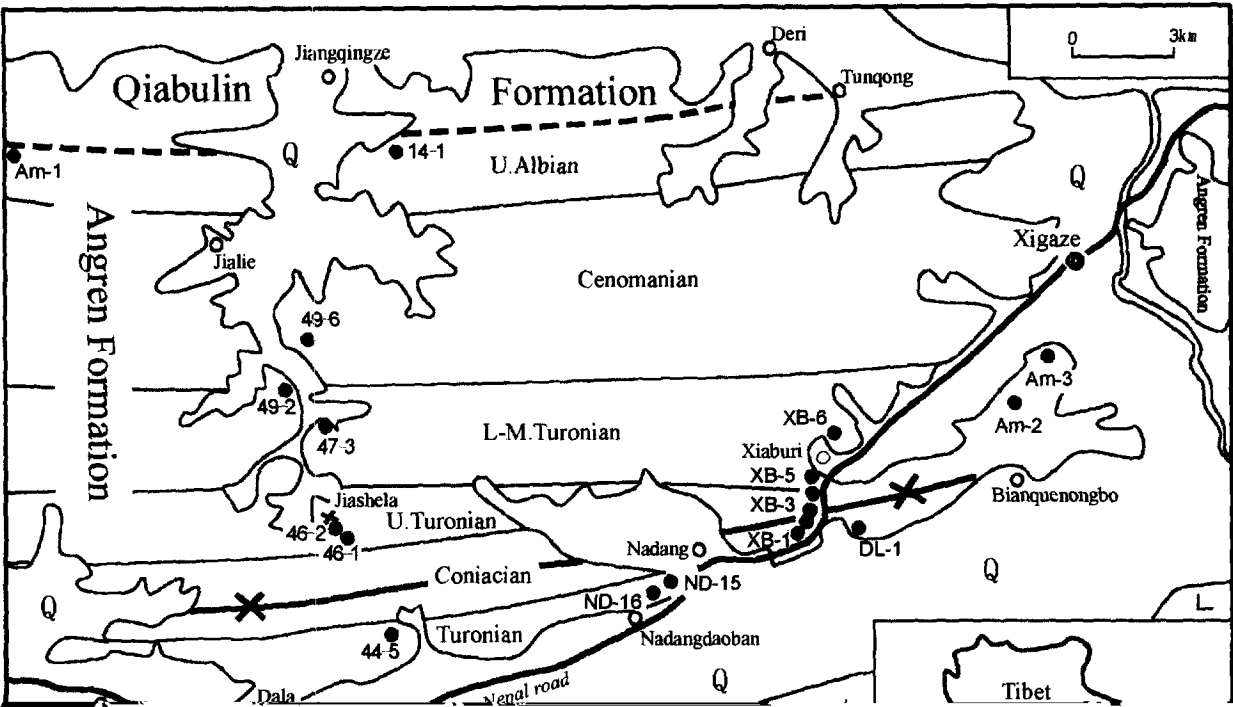
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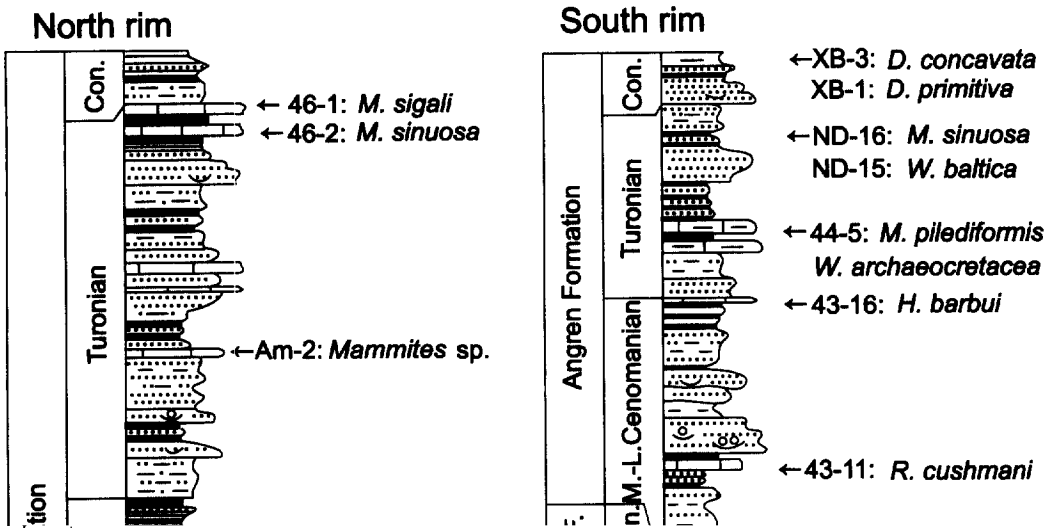
Abstract—The Xigaze Cretaceous forearc basin located between the Lhasa Block and Indus-Yarlung Suture was filled with flysch deposits during the mid-Cretaceous, as the Tethys oceanic crust was being subducted underneath Eurasia. The flysch comprises the Xigaze Group, dominated by the Agassiz Formation. Well-exposed and easily accessible fossiliferous strata of the Agassiz Formation are found in the Xigaze Forearc Basin, Tibet. The Agassiz Formation is a thick sequence of flysch deposits, which is characterized by a variety of fossiliferous strata. The fossils include a variety of marine organisms, including ammonites, bivalves, and brachiopods. The fossils are well-preserved and provide important information about the Cretaceous marine environment in the Xigaze Forearc Basin.

mation constrain its age to late Albian to late Coniacian. Radiometric dating of the ophiolite belt impinging on the southern edge of the forearc basin yielded an age of 120 ± 10 Ma, which implies that the deposition in the forearc basin occurred shortly after initiation of the subduction of the Neotethys. © 1998 Elsevier Science Ltd. All rights reserved

Introduction

Shixuan, 1974; Wu Haoruo *et al.*, 1977). This group





ND16: Greyish yellow marl contain a few foraminifera such as *Marginotruncana sinuosa* Porthault.

Section south from Xiaburi

Formation and therefore they provide better biostratigraphic control.

The oldest fossil assemblage in the Angren Formation is represented by sample 14-1. The corals

DL-1: Greyish yellow calcareous mudstone, yielding a few specimens of *Hedbergella delrioensis* (Carsey).

XB-1: Yellowish grey marl yielding many planktonic

ger foraminifera *Orbitolina* spp. is of Aptian-Albian age in this region. The ammonite *Neophlycticeras* (*N.*) *brottianum* which belongs to the *Mortoniceras inflatum* Zone (Wedderson and Diller 1995) occurs underneath

Marginotruncana sigali (Reichel), *Lenticulina* sp.

XB-2: Yellowish marl intercalated in black shale with few specimens of *Lenticulina* sp.

XB-3: Yellowish grey marl yielding abundant fora-

the sample 14-1 and is considered to be of an early late Albian age. Accordingly, the sample 14-1 is younger than early late Albian and its age is constrained by the occurrence of orbitolinids to the late



Fig. 3. SEM photomicrographs of Cretaceous foraminifera from Xigaze, Tibet. All specimens are deposited in the China University of Geosciences (Beijing). (1) *Hedbergella delrioensis* (Carsey) dorsal side; Sample ND15. $\times 160$. (2) *Hedbergella planispira* (Tappan) 2a. dorsal side; 2b. ventral side. sample 4311. $\times 160$. (3) *Hedbergella holmdelensis* Olsson 3a. dorsal side; 3b. ventral side. sample 4311, $\times 160$. (4) *Whiteinella inornata* (Bolli) 4a. dorsal side; 4b ventral side. Sample ND2790. (5) *Whiteinella holmdelensis* Olsson 5a. dorsal side; 5b. ventral side. Sample 445. $\times 80$. (6) *Whiteinella hol*

late Cenomanian *R. cushmani*–*H. barbui* assemblage; Turonian *Mammites* assemblage; middle Turonian *M. pillediformis*–*W. archaeocretacea* assemblage; late Turonian *M. sinuosa*–*M. sigali* assemblage; early Coniacian *D. primitiva* assemblage; and late Coniacian *D. concavata* assemblage.

of the Lhasa Block developed from late Albian to late Coniacian. The Aptian age of the Indus–Yarlung ophiolites (Gopel *et al.* 1984) indicates that the development of the forearc basin began early after initiation of the subduction of Neotethys. The new biostratigraphic data constrains the initiation of subduction of the Neotethys to early Albian.

Discussion

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